Do we need various assumptions to get a good FCN?

-- A new multiple layer spectral method

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- Methodology:
 - Multiple layer spectral method: Finite Element Method
 - Linear Operator Method
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Study of Free-core-nutation (FCN)

- FCN is a normal mode of the earth as the rotating axes of the FOC and of the mantle don't coincide.
- a key parameter & key question to be answered:
 - The calculated period of FCN from traditional theory differs largely from the high-precision obs.
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 - FCN reflects (depends on) the physics of the FOC, mantle & CN
 - FCN influences strongly the -1yr. nutation due to its resonance.



No fluid OC? → No FCN !

VLBI	430±1
SG(GGP)	430±5
theoretical calculated	458~470

- extra flattening at CMB: +5%: too big to be consistent with the overall near-HE shape of the Earth as a whole.
- magnetic/ viscous/topographic couplings @CMB,
- 2nd order (ε²) terms effect?
- etc.

Contributions of EMC @CMB to nutation (µas)



Mathews & Guo (2005); Buffett(2011): + viscous

Do /can we really need these various unproved assumptions to get a good FCN ?



Multiple layer spectral method (MLSM)

Linear Operator Method (LOM)

Finite Element Method (FEM)

- Traditional approach solves one order ellipsoid only.
- FEM can solve more complex models.



 $r = 7 - \cos(4\phi) * \sin^4 \theta$

 $r = 6 - 0.5(5 * \cos^3 \theta - 3\cos \theta)$

Main Idea of FEM

Boundary Surface could be described as:

$$r = r_0 + \sum_{n,m} \kappa_n^m Y_n^m(\theta,\phi)$$

Let's consider how to solve the dynamic equation:

$$\rho D_t^2 \overrightarrow{u} + 2\rho \overrightarrow{\Omega}_0 \times D_t \overrightarrow{u} = -\rho \overrightarrow{\Omega}_0 \times (\overrightarrow{\Omega}_0 \times \overrightarrow{u}) + \nabla \cdot \overleftrightarrow{S}^e - \nabla (\gamma \nabla \cdot \overrightarrow{u}) \\ -\rho \nabla \phi_1 - \rho \overrightarrow{u} \cdot \nabla \nabla \phi + \nabla \cdot [\gamma (\nabla \overrightarrow{u})^T]$$

Traditional Approach: Equivalent Spherical Domain



Traditional Approach: Equations of Simple Example

Governing Equations:

$$\rho \frac{\partial^2 \vec{u}}{\partial t^2} = \nabla \cdot \overleftrightarrow{S} - \rho (\nabla \cdot \vec{u}) \vec{g_0} + \rho \nabla \vec{u} \cdot \vec{g_0} + \rho \vec{u} \cdot \nabla \cdot \vec{g_0} + \rho \nabla V_1$$

$$\overleftrightarrow{S} = \lambda (\nabla \cdot \vec{u}) \overleftrightarrow{I} + \mu [\nabla \vec{u} + (\nabla \vec{u})^T]$$

$$\nabla^2 V_1 = 4\pi G \nabla \cdot (\rho \vec{u})$$

Variables:

$$\vec{u}(\vec{r},t) = \sum_{n=0}^{\infty} \sum_{m=-n}^{n} (\vec{S}_{nm}(\vec{r}) + \vec{T}_{nm}(\vec{r}))e^{i\omega}$$
$$V_1(\vec{r},t) = \sum_{n=0}^{\infty} \sum_{m=-n}^{n} V_{1nm}(r)y_{nm}(\theta,\phi)e^{i\omega t}$$

and

$$\overrightarrow{S_{nm}} = [U_{nm}(r)P_{nm}(\cos\theta)\hat{r} + V_{nm}(r)\frac{d}{d\theta}P_{nm}(\cos\theta)\hat{\theta} + V_{nm}(r)\frac{im}{\sin\theta}P_{nm}(\cos\theta)\hat{\phi}]e^{im\phi}$$
$$\overrightarrow{T_{nm}} = [W_{nm}(r)\frac{im}{\sin\theta}P_{nm}(\cos\theta)\hat{\theta} - W_{nm}(r)\frac{d}{d\theta}P_{nm}(\cos\theta)\hat{\phi}]e^{im\phi}$$

Traditional Approach: Direct numerical integration approach

Variables

Ordinary Differential Equations

$$y_1^n = U_{nm}$$

$$y_2^n = \lambda \triangle_{nm} + 2\mu \frac{dU_{nm}}{dr}$$

$$y_3^n = V_{nm}$$

$$y_4^n = \mu \left(\frac{dV_{nm}}{dr} + \frac{U_{nm} - V_{nm}}{r}\right)$$

$$y_5^n = V_{1nm}$$

$$y_6^n = \frac{dV_{1nm}}{dr} - 4\pi G\rho U_{nm}$$

$$y_7^n = iW_{nm}$$

$$y_8^n = i\mu \left(\frac{dW_{nm}}{dr} - \frac{W_{nm}}{r}\right)$$

$$\begin{cases} \frac{dy_1^n}{dr} = -\frac{2\lambda}{(\lambda+2\mu)r}y_1^n + \frac{1}{\lambda+2\mu}y_2^n + \frac{n(n+1)\lambda}{(\lambda+2\mu)r}y_3^n \\ \frac{dy_2^n}{dr} = \left[-\rho\omega^2 - \frac{4\rho g}{r} + \frac{4\mu(3\lambda+2\mu)}{(\lambda+2\mu)r^2}\right]y_1^n - \frac{4\mu}{(\lambda+2\mu)r}y_2^n + \\ \left[\frac{n(n+1)\rho g}{r} - \frac{2n(n+1)\mu(3\lambda+2\mu)}{(\lambda+2\mu)r^2}\right]y_3^n + \frac{n(n+1)}{r}y_4^n - \rho y_6^n \\ \frac{dy_3^n}{dr} = -\frac{1}{r}y_1^n + \frac{1}{r}y_3^n + \frac{1}{\mu}y_4^n \\ \frac{dy_4^n}{dr} = \left[\frac{\rho g}{r} - \frac{2\mu(3\lambda+2\mu)}{(\lambda+2\mu)r^2}\right]y_1^n - \frac{\lambda}{(\lambda+2\mu)r}y_2^n + \\ \left\{-\rho\omega^2 + \frac{2\mu}{(\lambda+2\mu)r^2}\right](2n^2 + 2n - 1)\lambda + 2(n^2 + n - 1)\mu]\right\}y_3^n \\ -\frac{3}{r}y_4^n - \frac{\rho}{r}y_5^n \\ \frac{dy_6^n}{dr} = -\frac{4\pi G\rho y_1^n + y_6^n}{r} \\ \frac{dy_6^n}{dr} = -\frac{4\pi G\rho n(n+1)}{r}y_3^n + \frac{n(n+1)}{r^2}y_5^n - \frac{2}{r}y_6^n \end{cases}$$

Traditional Approach: Direct numerical integration approach

- Give a possible period and ODEs' initial values
- Integrate ODEs from center to surface
- Check the integrated values at surface so as to determine whether the possible period is veritable



Traditional Approach Problem: the More Complex Figures ? ×



 $r = 1.5 + 1.3 \sin \theta$

 $r = 2.5 + 2\cos^4\theta$

 $r = 8 - cos^2 heta \qquad r = 7 - cos(4\phi) * sin^4 heta \qquad r = 6 - 0.5(5 * cos^3 heta - 3 cos heta)$

FEM: The Whole Domain is separated into several subdomain

eg, Earth could be separated into 3 subdomains:

- Solid inner core
- Fluid outer core
- Solid mantle (crust)



FEM: Express the Equation in Each Subdomain

The inner and the outer Boundary Surfaces could be described as:

$$r = R_{in} + \sum_{n,m} \xi_n^m Y_n^m(\theta,\phi)$$
$$r = R_{out} + \sum_{n,m} \Xi_n^m Y_n^m(\theta,\phi)$$

Let's consider how to solve the dynamic equation:

$$\rho D_t^2 \overrightarrow{u} + 2\rho \overrightarrow{\Omega}_0 \times D_t \overrightarrow{u} = -\rho \overrightarrow{\Omega}_0 \times (\overrightarrow{\Omega}_0 \times \overrightarrow{u}) + \nabla \cdot \overleftrightarrow{S}^e - \nabla (\gamma \nabla \cdot \overrightarrow{u}) \\ -\rho \nabla \phi_1 - \rho \overrightarrow{u} \cdot \nabla \nabla \phi + \nabla \cdot [\gamma (\nabla \overrightarrow{u})^T]$$

FEM -> multiple layer spectral method (MLSM)



multiple layer spectral method (MLSM): Integrate governing eqs. in each subdomain



We use MLSM instead of direct numerical integration approach. Variables are expanded in basis function series

MLSM: Galerkin Method

eg, the (2,0) spheroidal displacement field is represented as:



eg., trial vector(function) of (2,0) spheroidal is:

 $\phi_k(r) \cdot \nabla Y_2^0(\theta, \phi)$

MLSM: Boundary Conditions

Boundary surface could be described as:

$$r = r_0 + \sum_{n,m} \kappa_n^m Y_n^m(\theta,\phi)$$

The Normal vector of boundary surface is:

$$\overrightarrow{n} = \frac{\partial \overrightarrow{r}}{\partial \theta} \times \frac{\partial \overrightarrow{r}}{\partial \phi}$$

MLSM: Boundary Conditions turn into surface integral

for
$$\delta \vec{u} = \vec{u}^+ - \vec{u}^- = 0$$

$$\oint_{\theta=0}^{\pi} \sigma_{(i,j,k)} \hat{n} \cdot \delta \vec{u} dS$$

$$= \int_{\theta=0}^{\pi} \int_{\phi=0}^{2\pi} \sigma_{(i,j,k)} \{ [\hat{r} - \nabla r(\theta, \phi)] \cdot \delta \vec{u} \} [r(\theta, \phi)]^2 \sin \theta d\theta d\phi = 0$$

For scalar continuation condition $\delta a = a^+ - a^- = 0$

$$\iint_{\theta=0}^{\pi} \left| \frac{\partial \vec{r}}{\partial \theta} \times \frac{\partial \vec{r}}{\partial \phi} \right| \sin \theta * \delta a * \left| \frac{\partial \vec{r}}{\partial \theta} \times \frac{\partial \vec{r}}{\partial \phi} \right| dS$$
$$= \int_{\theta=0}^{\pi} \int_{\phi=0}^{2\pi} \left| \frac{\partial \vec{r}}{\partial \theta} \times \frac{\partial \vec{r}}{\partial \phi} \right|^{2} * \delta a * \sin \theta * d\theta d\phi = 0$$

MLSM: Combine All Equations

- Combine these equations so as to build a matrix:
 - Volume integral of governing equations in each subdomain
 - Surface integral of boundary conditions between 2 adjoint subdomain
 - Free surface boundary condition
- No need for the initial value at center. As center condition just require parameters to be reasonable which is the absence of r or r² basis terms.

MLSM: Search the Period

 Pick an period and compute the condition number of the matrix.

Linear Operator Method (LOM)

- Why Use Linear Operator method?
 - Generalized Spherical Harmonics(GSH) are a little bit abstruse. It needs knowledge of group theory and representation theory.
 - Boundary conditions could be easily solved.

Linear Operator Method (LOM)

 Equations is based on spherical harmonics (SH) with unified form:

$$Y_n^m(\theta,\phi) = \sum_{s=0}^{[(n-|m|)/2]} c_{10}(n,m,s)(\cos\theta)^{n-2s-|m|} (\sin\theta e^{\zeta(m)i\phi})^{|m|}$$

• Each SH can be built up by 3 atoms:

 $\cos\theta, \sin\theta e^{i\phi}, \sin\theta e^{-i\phi}$

 If we know 3 atoms' actions on each other, all computation about SH are obtained.

LOM Example: Product of two Spherical Harmonics

As we know:

 $\cos\theta * Y_n^m(\theta,\phi) = d_1(n,m) * Y_{n+1}^m(\theta,\phi) + d_2(n,m) * Y_{n-1}^m(\theta,\phi)$ $\sin\theta e^{i\phi} * Y_n^m(\theta,\phi) = d_{10}(n,m)Y_{n+1}^{m+1}(\theta,\phi) + d_{11}(n,m)Y_{n-1}^{m+1}(\theta,\phi)$ $\sin\theta e^{-i\phi} * Y_n^m(\theta,\phi) = d_{12}(n,m)Y_{n+1}^{m-1}(\theta,\phi) + d_{13}(n,m)Y_{n-1}^{m-1}(\theta,\phi)$

$$Y_n^m(\theta,\phi) = \sum_{s=0}^{[(n-|m|)/2]} c_{10}(n,m,s)(\cos\theta)^{n-2s-|m|} (\sin\theta e^{\zeta(m)i\phi})^{|m|}$$

We can get

• By

 $Y_n^m(\theta,\phi) * Y_i^j(\theta,\phi)$

LOM Example: Product of a SH and a VSH

• As we know: $\cos \theta * \overrightarrow{S}_{n}^{m}(\theta, \phi) = d_{7}(n, m) * \overrightarrow{S}_{n+1}^{m}(\theta, \phi) + d_{8}(n, m) * \overrightarrow{S}_{n-1}^{m}(\theta, \phi)$ $+ d_{9}(n, m) * \overrightarrow{T}_{n}^{m}(\theta, \phi)$ $\sin \theta e^{i\phi} * \overrightarrow{S}_{n}^{m}(\theta, \phi) = d_{14}(n, m) * \overrightarrow{S}_{n+1}^{m+1}(\theta, \phi) + d_{15}(n, m) * \overrightarrow{S}_{n-1}^{m+1}(\theta, \phi)$ $+ d_{16}(n, m) * \overrightarrow{T}_{n}^{m+1}(\theta, \phi)$

$$\sin \theta e^{-i\phi} * \overrightarrow{S}_n^m(\theta, \phi) = d_{17}(n, m) * \overrightarrow{S}_{n+1}^{m-1}(\theta, \phi) + d_{18}(n, m) * \overrightarrow{S}_{n-1}^{m-1}(\theta, \phi)$$
$$+ d_{19}(n, m) * \overrightarrow{T}_n^{m-1}(\theta, \phi)$$

• We can get

$$Y_a^b(\theta,\phi) * \nabla_1 Y_n^m(\theta,\phi)$$

LOM Example: Dot-product of two VSH

Vector Spheroidal harmonics (VSH) can be written in following form:

 $\overline{\nabla_1 Y_n^m(\theta,\phi)} = H_0[Y_n^m(\theta,\phi)] * \overline{\nabla_1 \cos \theta} + H_1[Y_n^m(\theta,\phi)] * \overline{\nabla_1 (\sin \theta e^{\zeta(m)i\phi})}$

• H_0 and H_1 are combination of SH:

$$H_0[Y_n^m(\theta,\phi)] = \sum_{s=0}^{[(n-|m|)/2]} (n-2s-|m|) * c_{10}(n,m,s)(\sin\theta e^{\zeta(m)i\phi})^{|m|} \\ * (\cos\theta)^{n-2s-|m|-1} \\ H_1[Y_n^m(\theta,\phi)] = \sum_{s=0}^{[(n-|m|)/2]} |m| * c_{10}(n,m,s)(\cos\theta)^{n-2s-|m|} \\ * (\sin\theta e^{\zeta(m)i\phi})^{|m|-1}$$

LOM Example: Dot-product of two VSH

• If we get:

 $\nabla_1 \cos \theta \cdot \nabla_1 Y_a^b(\theta, \phi)$ $\nabla_1 (\sin \theta e^{i\phi}) \cdot \nabla_1 Y_a^b(\theta, \phi)$ $\nabla_1 (\sin \theta e^{-i\phi}) \cdot \nabla_1 Y_a^b(\theta, \phi)$

• We can finally get:

 $\nabla_1 Y_n^m(\theta,\phi) \cdot \nabla_1 Y_a^b(\theta,\phi)$

LOM

- Use this method, we get
 - product of
 - 2 SH
 - a SH & a VSH
 - dot product of 2 VSH
 - cross product of 2 VSH
 - gradient of a SH
 - curl of a VSH
 - divergence of a VSH

LOM: Tensors

 It is difficult to represent the stress tensor in a stand-alone form. But in the equation it only needs the divergence of the tensor, while in boundary conditions it only needs the dotproduct of the normal vector and the tensor, and these two terms can be represented by the LOM.

Our Preliminary Earth Model



• PREM

- 1 order ellipticity
- · 3 layers
 - without ocean
 - 1 layer for solid inner core
 - 1 layer for fluid outer core
 - 1 layer for mantle and crust (while 10 layers of parameters)

Validation: Tilt-Over-Mode

 The displacement field is truncated as:

$$\overrightarrow{T}_{1}^{1} + \overrightarrow{S}_{2}^{1} + \overrightarrow{T}_{3}^{1}$$

Each term is expanded in polynomial series, in each subdomain

$$\overrightarrow{T}_{1}^{1} = \sum_{i=0}^{Max-order} c_{i} * r^{i} * [\widehat{r} \times \nabla_{1}Y_{1}^{1}(\theta, \phi)]$$



ΤΟΜ

Order of Polynomial in trial function

FCN Result

 The displacement field is truncated as:

 $\overrightarrow{T}_{1}^{1} + \overrightarrow{S}_{2}^{1} + \overrightarrow{T}_{3}^{1}$

- FCN is very sensitive to the ellipticity at CMB. Our value is equivalent to the most authors': 2.54656*10⁻³.
- Result converges when polynomial order i_{max} >= 4.

VLBI	430±1
SG(GGP)	430±5
Theory calculated	458~470
Our Approach	435±3

Discussion : Why MLSM is Better?

 avoiding derivatives of some parameters which are not precise in earth model.

$$\lambda(P) = \lambda_0(P) + \frac{2}{3}\epsilon(P)p\frac{\partial\lambda_0(P)}{\partial p}P_2(\cos\theta)$$
$$\int_{r} \partial_r(\rho,\lambda,\mu)dr \Longrightarrow(\rho,\lambda,\mu)$$

$$\partial_r(\rho,\lambda,\mu)$$

MLSM focuses on global characteristics •

$$\int_{\theta=0}^{\pi} \int_{\phi=0}^{2\pi} \int_{R_{in}+\sum_{n}^{m} \xi_{n}^{m} Y_{n}^{m}(\theta,\phi)} \overrightarrow{\Lambda}_{(i,j,k)} \cdot \overrightarrow{eq}_{v} r^{2} \sin \theta dr d\theta d\phi = 0$$

Discussion & Next Step?

- Clairaut coordinates (Rochester et al. ,2014): =458
- 2^{nd} order (ε^2) terms effect ?
- truncated coupling chain: $\vec{T}_1^1 + \vec{S}_2^1 + \vec{T}_3^1 + \dots$?

Rotational modes of Jupiter, Saturn & exoplanet

Thanks!

Study of Free-core-nutation (FCN)

- FCN is a normal mode of the earth as the rotating axes of the FOC and of the mantle don't coincident.
- FCN depends on the physics of the FOC, mantle & core-mantle-boundary. It influences strongly the retro-annual nutation due to its resonance, so it is a key parameter & interesting topic.
- The calculated period of FCN from traditional theory differs largely from the high-precision observation.
- We developed an integrated Galerkin method and spectral element method that can study any antisymmetric earth without GSH.
- These methods are applied on the computation of FCN period from PREM earth without any assumption (eg., extra flattening at CMB, magnetic/ viscous/topographic couplings at CMB, etc.).
- Our result is 435 sid. day !



(Zhang & Huang, 2014a,b, c)

This work

Mantle

435

FOC